

1    Orientation and intensity of maximum response spectral  
2            ordinates during the December 20, 2022  $M_w$  6.4  
3            Ferndale, California earthquake

4            Nathan Girmay<sup>1\*</sup>, Eduardo Miranda<sup>1</sup>, Alan Poulos<sup>1</sup>

5            <sup>1</sup>Department of Civil and Environmental Engineering, Stanford University,  
6            Stanford, California, U.S.A

7            \* Corresponding author. *Email*: ngirmay3@stanford.edu

8    **Highlights**

- 9    • Orientation of the maximum response occurred close to the transverse orientation.  
10    • This is consistent with recent observations from other strike-slip earthquakes.  
11    • Spectral response at transverse orientation, on average, was close to RotD100.  
12    • This was observed for oscillators with periods ranging from 0.1 s to 10 s.

13    **Keywords**

14    Ground motion directionality; Transverse orientation; Ground motion polarization; Response spectral  
15    ordinates; Ground motion intensity.

16 **Abstract**

17           The level of horizontal ground motion intensity can have large variations from one orientation to  
18 another. This study examines the orientation of the maximum response of 5%-damped linear oscillators and  
19 its geographical distribution for ground motions recorded during the December 20, 2022, Ferndale, CA  
20 earthquake. The orientation of the maximum intensity is found to generally occur close to the transverse  
21 orientation for epicentral distances greater than 30 km. At epicentral distances less than 30 km, the angular  
22 difference between the orientation of the maximum intensity and transverse orientation does not show a  
23 clear predominant orientation. The angular difference with respect to the transverse orientation is found to  
24 be strongly correlated with the level of intensity. The ground motion intensity at the transverse orientation  
25 was, on average, between 85% to 95% of the maximum intensity and was 1.0 to 1.26 times greater than the  
26 RotD50 intensity (i.e., the median intensity from all horizontal orientations). The level of polarization  
27 increases with the period of the oscillation and ground motions remain fairly polarized up to 350 km from  
28 the epicenter. The results support a case for future incorporation of directionality effects, and in particular,  
29 the use of orientation relative to the transverse orientation, when estimating response spectral ordinates  
30 from strike-slip earthquakes.

## 31 **1. Introduction**

32 In earthquake-resistant design, the most common measure of ground motion intensity is the 5%-  
33 damped response spectral ordinate, which represents the peak response of a single-degree-of-freedom  
34 linear-elastic oscillator with 5% damping subjected to a given ground motion. It has long been known that  
35 horizontal ground motion intensity, including the commonly used 5%-damped response spectral ordinate,  
36 varies with changes in orientation (i.e., changes in the azimuth considered at the given station) in what is  
37 referred to as directionality. Although horizontal earthquake ground motion intensity, in general, is  
38 recorded in two perpendicular orientations and exhibits different amplitudes in different horizontal  
39 orientations, only a single measure of ground motion intensity is typically used in current ground motion  
40 models (GMMs), essentially ignoring directionality effects. There have been several approaches for  
41 obtaining a single measure of intensity to represent the horizontal ground motion intensity that occurred at  
42 a site. The four most common methods that have been used in the past are: (1) combining the two as-  
43 recorded horizontal intensities such as by using their geometric mean (e.g., Joyner and Boore, 1982; Boore  
44 et al., 1997; Abrahamson and Silva, 1997; Beyer and Bommer, 2006); (2) selecting one of the two as-  
45 recorded horizontal intensities, what has typically been referred to as the arbitrary component (e.g., Boore  
46 et al., 1997); (3) using a horizontal intensity that is independent of the orientation of horizontal sensors  
47 (e.g., Boore et al., 2006; Boore et al., 2010); and (4) selecting the intensity occurring at a specific orientation  
48 with respect to the strike of the rupture such as strike-normal and strike-parallel components (e.g.,  
49 Somerville et al., 1997).

50 Recent studies have shown that directionality effects can be significant. In other words, at a given  
51 site, the level of intensity can have large variations from one horizontal orientation to another (Hong and  
52 Goda, 2007; Shahi and Baker, 2014; Poulos and Miranda, 2022b). For example, using the NGA-West2  
53 ground motion database, Poulos et al. (2022) found that 5%-damped response spectral ordinates at the  
54 maximum intensity are, on average, 1.3 to 2.0 times higher than the response spectral ordinate in the  
55 perpendicular direction depending on the period of the oscillator. However, for many years, directionality  
56 was not taken into account for design in the United States (U.S.) and it is still neglected in most countries.  
57 In the U.S., directionality was considered indirectly, starting with the 2010 version of ASCE 7, by using  
58 approximate amplification factors to estimate the intensity at the orientation of the maximum intensity  
59 (RotD100) from the median intensity from all orientations (RotD50). Since most GMMs do not use  
60 RotD100, ASCE 7 uses approximate period-dependent amplification factors to construct the design  
61 spectrum (ASCE, 2010; Poulos and Miranda, 2022a). Nonetheless, the factors used for the short-period  
62 response spectral region (0.2 s) in ASCE 7-10 underestimated the required amplification (Shahi and Baker,

63 2014). This has now been amended, with ASCE 7-22 incorporating improved period-dependent  
64 amplification factors at 22 different periods instead of only at 0.2 s and 1.0 s (ASCE, 2022).

65 Except for near-fault sites (i.e., source-to-site distances shorter than 5 to 15 km), the orientation of  
66 the maximum intensity is currently considered to be uniformly distributed (Huang et al., 2008; NEHRP  
67 Consultants Joint Venture, 2011; Shahi and Baker, 2014). Accordingly, current seismic design standards in  
68 the U.S. emphasize the lack of predominant orientation in sites that are not near-fault (ASCE, 2010; ASCE,  
69 2016; ASCE, 2022; PEER, 2017).

70 Whilst there have been several investigations on the directionality of response spectral ordinates,  
71 practically all of them have focused only on whether the orientation of the maximum response occurs at or  
72 close to the strike-normal orientation. In other words, they were studying if the orientation of the maximum  
73 intensity was approximately the same in all recording stations and how close to the strike normal it was.  
74 Recently, Poulos and Miranda (2023) investigated the orientation of the maximum intensity of ground  
75 motion records from shallow crustal earthquakes in the NGA-West2 database (Ancheta et al., 2014),  
76 studying them separately for each event for a total of 1,966 and 2,226 ground motions recorded in strike-  
77 slip and reverse faulting earthquakes, respectively. Using these recorded ground motions, they showed that  
78 for strike-slip events, the orientation of the maximum intensity tended to be relatively close to the transverse  
79 orientation (i.e., an orientation perpendicular to the line segment connecting the epicenter to the recording  
80 station). Furthermore, they showed that the tendency for the orientation of the maximum intensity to be  
81 closer to the transverse orientation increases with increasing period of the oscillator. This observation by  
82 Poulos and Miranda (2023) is important as it suggests that the orientation of the maximum response of  
83 future earthquakes can be estimated from the position of the epicenter relative to the site of interest, thus  
84 improving the ability to estimate ground motion intensity at different orientations.

85 The December 20<sup>th</sup>, 2022,  $M_w$  6.4 Ferndale earthquake is a well-recorded event with a strike-slip  
86 faulting style (with a rake angle of  $7^\circ$  (USGS, 2022)) that provides an important opportunity to  
87 independently evaluate the observations by Poulos and Miranda (2023). It also provides an opportunity to  
88 evaluate the amplitude of response spectral ordinates in the transverse orientation relative to the RotD100  
89 intensity (which is currently used by design standards in the U.S.) and the RotD50 intensity (which is  
90 commonly used in recent GMMs). This information could then be used in combination with recent  
91 probabilistic directionality models (Poulos and Miranda, 2022a; 2022b) for the development of correction  
92 factors to existing RotD50 GMM to then estimate levels of ground motion intensity at specific orientations.

93 This work examines the orientation of the maximum response for 5%-damped linear-elastic  
94 oscillators subjected to ground motions recorded in the December 20, 2022, Ferndale earthquake in northern  
95 California. Additionally, it examines several aspects of the directionality of ground motions recorded in

96 this event that were not studied by Poulos and Miranda (2023). In particular, the following aspects are  
97 studied: (a) geographical distribution of the level of polarization of ground motions, which in this paper  
98 refers to the amount of directionality, and its variation with changes in the period of vibration for linear-  
99 elastic oscillators; (b) geographic distribution of the orientation of the maximum response spectral ordinate  
100 for linear-elastic oscillators with different periods of vibration and the effect of source-to-site distance on  
101 this orientation; (c) level of amplitude in response spectral ordinates in the transverse and radial orientations  
102 relative to the maximum intensity (RotD100); (d) level of amplitude in response spectral ordinates in the  
103 transverse and radial orientations relative to RotD50 intensities; (e) probability of exceeding RotD50  
104 intensities in transverse and radial orientations.

## 105 **2. Selected earthquake ground motions**

106 In this paper, records from the 2022  $M_w$  6.4 Ferndale earthquake are used to study the directionality  
107 of horizontal ground motion. The earthquake occurred on December 20, 2022, at 10:34 (UTC) (2:34 pm  
108 local time), with an epicenter located close to the coast of northern California, approximately 15 km  
109 southwest of Ferndale and 30 km south of Eureka, California (USGS, 2022). Focal mechanism suggests  
110 strike-slip faulting (with a rake angle of  $7^\circ$ ) on a steeply dipping fault that strikes west-southwest. The  
111 epicenter is in the most seismically active region in California due to the intersection of three tectonic plates  
112 (i.e., the Pacific Plate, the North American Plate, and the Juan de Fuca Plate) known as the Mendocino  
113 Triple Junction. At the intersection of the three plates is the crossing of the northern tip of the San Andreas  
114 Fault, the Mendocino Fracture Zone, and the Cascadia Subduction Zone (Temblor, 2022). This region  
115 experienced a sequence of overlapping earthquakes of  $M_w$  6.2 and 5.7 almost a year before the 2022  
116 earthquake. In the past century, there have been at least 40 other earthquakes with  $M_w \geq 6$ , including six  
117 with  $M_w \geq 7$ , within roughly 250 km of the 2022 earthquake (USGS, 2022). For a detailed description of  
118 the seismicity of this region of California, the reader is referred to Bakum (2000). The earthquake resulted  
119 in widely distributed power loss and damage to numerous buildings in Humboldt County, with over 30  
120 residential and commercial buildings found structurally unsafe (Petri and Lin, 2022).

121 For this study, ground motion data, as processed by California Geological Survey  
122 (CGS)/California Strong Motion Instrumentation Program (CSMIP) and United States Geologic  
123 Survey/National Strong Motion Project (NSMP) was obtained from the Center for Engineering Strong  
124 Motion Data (CESMD). The data processing performed by CGS/CSMIP or USGS/NSMP includes  
125 instrument correction, baseline correction, and band-pass filtering using acausal filters (CESMD).  
126 Processed records (i.e., Volume 2) included in the study were selected according to the following criteria:  
127 (1) the record must be available in both horizontal components at a given station; (2) the orientation and  
128 polarity of the horizontal sensors at the recording station must be available; (3) the recording station must

129 be a free-field station or acceptably represent free-field conditions; (4) the corner frequency of the low-cut  
130 (high-pass) filter at both horizontal components must be shorter than 0.08 Hz; and (5) at least one of the  
131 two recorded horizontal components must have a peak ground velocity (PGV) greater than or equal to 1  
132 cm/s.

133         Since previous studies have found that polarization increases with period (e.g., Poulos and Miranda,  
134 2022b; 2023), it is of interest to consider records that allow the reliable calculation of response spectral  
135 ordinates at relatively long periods. Therefore, the fourth criterion ensures that the records have maximum  
136 usable periods equal or longer than 10 s, making them suitable to study directionality for oscillators with  
137 periods up to 10 s (Boore, 2004). The minimum level of ground motion intensity in the fifth criterion was  
138 selected as an easy-to-use criterion to achieve a strong signal-to-noise ratio up to periods of 10 s, again to  
139 be able to carefully examine directionality for oscillators with periods up to 10 s. Based on the selection  
140 criteria outlined above, 70 ground motion records were found to be usable in this study for oscillators with  
141 natural periods ranging between 0.1 s and 8.0 s, whilst 55 ground motion records were found usable for  
142 oscillators with periods ranging between 8.0 s and 10 s.

### 143 **3. Orientation of maximum ground motion intensity**

144         For a given event, a linear-elastic oscillator within a horizontal plane can exhibit notable  
145 directionality and polarization (i.e., significantly larger response in certain orientations than in others). A  
146 simple way to visualize the response of an oscillator within a plane is to cross-plot the relative displacement  
147 of an oscillator when subjected to the two recorded horizontal components, resulting in a vectorial  
148 representation of displacement in the plane, in what is called a hodograph or particle motion trace. Figure  
149 1 shows the relative displacement hodographs of 5% damped linear elastic oscillators having periods of  
150 vibration of 10 s and 5 s when subjected to ground motions recorded during the 2022  $M_w$  6.4 Ferndale  
151 earthquake. The figure also shows the spatial distribution of the stations considered. In this figure, each  
152 hodograph has been normalized by its maximum spectral response (RotD100), which corresponds to the  
153 point farthest away from the origin (i.e., resting point) of the hodograph. Hence, all hodographs in this  
154 figure are shown with the same peak amplitude in order to better appreciate the directionality at each  
155 recording station. From this figure, it is apparent that stations close to each other tend to have response  
156 orientations that are similar to each other. This observation is only partly evident for recording stations that  
157 are near the epicenter (see the closeups on the right-hand side of Figure 1) but generally becomes more  
158 apparent at stations located further from the epicenter. Nonetheless, the hodographs visually demonstrate  
159 that the oscillator response at most stations is fairly polarized for this event, even at distances far from the  
160 source, with clearly noticeable directions of predominant response. It is to be noted that directionality is not  
161 necessarily related to forward directivity in the near-source region. As identified by Boore and Akkar

162 (2003), the intensity in two orthogonal orientations at a given site can be notably different, suggesting  
163 ground motion can be highly polarized even at distances very far from the source.

164 There are several ways in which the level of polarization in horizontal ground motion can be  
165 quantified. One of those is by using the ratio in intensities between the transverse and radial orientations.  
166 Figure 2 shows histograms of the ratio between ground motion intensities, as measured by response spectral  
167 ordinates, in the transverse orientation,  $S_{aT}$ , and the intensity in the radial orientation,  $S_{aR}$ , for periods of 3,  
168 5, 7, and 10 s. It can be seen that these ratios are typically larger than one, with mean values ranging from  
169 1.68 to 2.54 depending on the period of vibration of the oscillator, and that these ratios for some records at  
170 certain periods can be larger than five. Figure 3 shows the influence of the period of vibration on the level  
171 of polarization as measured by the mean ratio between the ground motion intensity at the transverse  
172 orientation and the intensity at the radial orientation. From this figure, it is clear that the level of  
173 polarization, on average, tends to increase as the fundamental period of the oscillator gets longer. Also  
174 shown in Figure 3 is the interquartile range from which one may infer that for periods of vibration longer  
175 than 2 s, there is at least a 75% chance that the 5%-damped spectral ordinate will be larger in the transverse  
176 orientation than in the radial orientation.

177 Figure 4 shows the orientation of the maximum response at each recording station of 5% damped  
178 linear elastic oscillators for periods of 3, 5, 7, and 10 s. At each recording station, this orientation of the  
179 maximum response is indicated by short black lines. As done by Poulos and Miranda (2023), the angular  
180 difference between the orientation of the maximum intensity and the transverse orientation is defined as the  
181 variable  $\alpha \in [-90^\circ, 90^\circ]$ , where a positive value corresponds to counterclockwise orientations with respect  
182 to the transverse orientation. The absolute angular difference between the orientation of the maximum  
183 intensity and the transverse orientation is  $|\alpha| \in [0^\circ, 90^\circ]$ . Figure 4 depicts the values of  $|\alpha|$  at each station  
184 by the color inside the small circles, with blue indicating values closer to  $0^\circ$  (i.e., closer to the transverse  
185 orientation) and red indicating values closer to  $90^\circ$  (i.e., closer to the radial orientation). Evidently, across  
186 the four periods, the large majority of recording stations are colored blue, suggesting that the predominant  
187 orientation of the maximum intensity is much closer to the transverse orientation than to the radial  
188 orientation.

189 Histograms of  $|\alpha|$  for oscillators with periods of 3, 5, 7, and 10 s are shown in Figure 5. It is  
190 readily apparent that the probability distribution of  $|\alpha|$  is noticeably skewed to low values of  $|\alpha|$ , with  
191 most of the records having values of  $|\alpha|$  below  $45^\circ$ . Should the orientation of the maximum intensity with  
192 respect to the transverse orientation be equally likely in any orientation, the histograms would appear as  
193 uniform distributions, as indicated by the blue discontinuous line in each of these histograms, and have a  
194 mean  $|\alpha|$  of  $45^\circ$ . However, the mean  $|\alpha|$  for the four periods is significantly below  $45^\circ$ , indicating that

195 the orientations of the maximum intensity for this event are, on average, relatively close to the transverse  
196 orientations, consistent with recent observations by Poulos and Miranda (2023) for strike-slip events in the  
197 NGA-West2 ground motion database. This observation is true for all periods and becomes more apparent  
198 as the period of the oscillator increases. Figure 6 shows the fraction of recording stations where the  
199 orientation of the maximum intensity falls within  $\pm 25^\circ$  of the transverse orientation. According to previous  
200 studies by Poulos and Miranda (2022b), within this range of orientations, the spectral ordinate will be, on  
201 average, within 10% of the maximum spectral ordinate from all orientations (i.e., within 10% of the  
202 RotD100 intensity). Should the orientation of RotD100 be fully random, that is, equally likely to occur in  
203 any orientation, the expected fraction of recording stations that would be within  $\pm 25^\circ$  would be only  
204 approximately 28% ( $50^\circ/180^\circ$ ). However, as shown in this figure, for oscillators with a period of  
205 approximately 3 s, orientations of RotD100 occurred within  $\pm 25^\circ$  of the transverse orientation in more than  
206 50% of the recording stations. Moreover, at a period of 7.5 s, this fraction of recording stations with  
207 orientations of the maximum response being within  $25^\circ$  of the transverse orientation increases to  
208 approximately 80%, which is significantly above and nearly three times larger than the 28% that would  
209 occur if the orientation of the maximum response were to be fully random, suggesting that the orientation  
210 of the maximum intensity is not uniform and that it occurs at orientations close to the transverse orientation.

211 Figure 7 shows the relationship between  $|\alpha|$  and the ratio of intensity at the transverse orientation  
212 to the maximum intensity at each recorded station for a 10 s oscillator. From the figure, it is apparent that  
213 the spectral ordinate in the transverse orientation relative to the RotD100 intensity is strongly correlated  
214 with  $|\alpha|$ , closely following the cosine curve. In particular, when  $\alpha$  is within  $\pm 25^\circ$  of the transverse  
215 direction, the spectral ordinate in the transverse orientation is within 10% of the RotD100 intensity,  
216 suggesting that this range of orientations close to the transverse orientation is the one that will experience  
217 the strongest intensities. This figure indicates that the value of  $|\alpha|$  could be used as a relatively good  
218 estimator of the level of ground motion intensity (i.e., smaller alpha translates to a larger intensity and  
219 approaches RotD100 at transverse orientation, approximately following the cosine). More importantly, the  
220 figure suggests that the drop in intensity at the transverse orientation is not significant for a notable range  
221 of  $|\alpha|$ . The level of ground motion intensity at the transverse orientation is discussed in greater detail in  
222 the following section of this paper.

223 While the influence of source-to-site distance on the orientation of the maximum intensity has been  
224 investigated on numerous occasions, the focus of most investigations has been the orientation with respect  
225 to strike-normal orientation (e.g., Huang et al., 2008; Shahi and Baker, 2014). To get a better understanding  
226 of the influence of distance on the orientation of the maximum intensity with respect to the transverse  
227 orientation, angular distances between the transverse orientation and RotD100 were binned by epicentral



228 distance. Figure 8 shows the influence of epicentral distance on the histograms of  $|\alpha|$  for oscillator periods  
229 of 5, 7, and 10 s. Although the number of records in each bin is relatively small for epicentral distances less  
230 than 30 km, the figure suggests that, for this earthquake, within the near field,  $|\alpha|$  is approximately  
231 uniformly distributed with mean  $|\alpha|$  values closer to  $45^\circ$  across the three periods of vibration. These results  
232 suggest that the orientation of the maximum intensity for near-fault sites does not necessarily occur close  
233 to the transverse orientation, but that, at least for this earthquake, they do not appear to correspond to the  
234 strike-normal orientation either as identified by others. However, for epicentral distances greater than 30  
235 km, the histograms are clearly skewed toward small values of  $|\alpha|$  and mean values of  $|\alpha|$  are significantly  
236 below  $45^\circ$ , indicating that for this event, the orientation of the maximum intensity is, on average, close to  
237 the transverse orientation, which is consistent with recent findings by Poulos and Miranda (2023) for strike-  
238 slip earthquakes.

#### 239 **4. Intensity at transverse and radial orientations**

240 From the results presented in the previous section, it is clear that the maximum intensity is not  
241 equally likely to occur in all horizontal orientations. While the computed mean  $|\alpha|$  values are notably  
242 smaller than  $45^\circ$ , they are not always exactly  $0^\circ$ . Therefore, it is of interest to investigate the ground motion  
243 intensity in the transverse and radial orientation at each recorded station with respect to both RotD100 and  
244 RotD50.

245 Herein, an evaluation of the intensity in the transverse and radial orientation with respect to  
246 RotD100 is performed by computing the intensity ratios  $S_{aT}/S_{a_{RotD100}}$  and  $S_{aR}/S_{a_{RotD100}}$ . Since RotD100 is  
247 the maximum intensity within the horizontal plane, the maximum possible intensity at the transverse or  
248 radial orientations is the RotD100 intensity. Thus, both  $S_{aT}/S_{a_{RotD100}}$  and  $S_{aR}/S_{a_{RotD100}}$  can range between 0  
249 and 1. Figure 9 shows histograms of the ratio between the ground motion intensity at the transverse or radial  
250 orientation and RotD100 computed from ratios at all recording stations for periods of 3, 5, 7, and 10 s. As  
251 seen in the figure, the probability distribution of  $S_{aT}/S_{a_{RotD100}}$  ratios is heavily skewed towards large values  
252 for all four periods, with most ratios being relatively close to 1, indicating that the intensity at the transverse  
253 orientation is not much lower than the maximum intensity. For example, the intensity at the transverse  
254 orientation for an oscillator with a fundamental period of 3 s is, on average, 91% of the maximum intensity.  
255 The average of  $S_{aT}/S_{a_{RotD100}}$  ratios increases to 95% of RotD100 at 10 s. This contrasts significantly with  
256 the intensity at the radial orientation where  $S_{aR}/S_{a_{RotD100}}$  ratios exhibit a much larger variability and have  
257 mean values close to 0.5, that is, with intensities in the radial orientation that are on average approximately  
258 half of the maximum intensity, highlighting the importance of ground motion directionality.

259 A better understanding of the influence of oscillator period on  $Sa_T/Sa_{RotD100}$  and  $Sa_R/Sa_{RotD100}$  can  
 260 be gained by considering how the mean ratio for all stations varies with changes in the period of vibration  
 261 of the oscillator. Figure 10 shows the mean ratio between the ground motion intensity at the transverse or  
 262 radial orientations and the RotD100 intensity for  $T \in [0.1 \text{ s}, 10 \text{ s}]$ . In this figure, variable  $\theta \in \{0^\circ, 90^\circ\}$  is  
 263 defined as the angular distance between the transverse orientation and the orientation of interest (i.e.,  $Sa(\theta$   
 264  $= 0^\circ)$  corresponds to the intensity at the transverse orientation). It can be seen that, even for short periods,  
 265 the intensity at the transverse orientation is above 80% of the RotD100 intensity. Furthermore, as the  
 266 oscillator period increases, the mean intensity at the transverse orientation also tends to increase, and by 3  
 267 s the average intensity at the transverse orientation is above 90% of the RotD100 intensity. In contrast, the  
 268 intensity at the radial orientation exhibits strong reductions as the period increases. These results indicate  
 269 that the significant increase in polarization with increasing periods is due not so much to the increase in  
 270 intensity in the transverse orientation but to the reduction in intensity in the radial orientation. In this figure,  
 271 the shaded areas around the means indicate the interquartile range for each period. It is discernable that  
 272 there is significantly less variability in the  $Sa_T/Sa_{RotD100}$  ratio than in the  $Sa_R/Sa_{RotD100}$  ratio, with the band  
 273 around the mean for the transverse orientation getting narrower with increasing period. Results shown in  
 274 Figures 9 and 10 have important practical implications as they indicate that prediction of the exact  
 275 orientation of RotD100 is not required since, for a wide range of angles close to the transverse direction,  
 276 there is an intensity which is relatively close to the maximum intensity. The mean  $Sa_T/Sa_{RotD100}$  ratio for  
 277 this earthquake can be roughly approximated as

$$\frac{Sa_T}{Sa_{RotD100}} = 0.85 + \frac{T}{85} \quad (1)$$

279 where  $T$  is the period of vibration of the oscillator in seconds. A similar fit can be performed using a  
 280 larger database of ground motions from strike-slip events to obtain a more generalized equation to  
 281 estimate the mean  $Sa_T/Sa_{RotD100}$  for linear elastic oscillators.

282 As discussed in the introduction, most current GMMs use a single scalar measure of intensity in  
 283 the horizontal orientation. Currently, the most commonly used scalar as a measure of horizontal ground  
 284 motion intensity in GMMs is RotD50, which, for a given period of vibration, corresponds to the median  
 285 intensity from all non-redundant horizontal orientations. Figure 11 shows the fraction of recording stations  
 286 where the intensity at the transverse or the radial orientation exceeded the RotD50 intensity for periods of  
 287 vibration ranging from 0.1 to 10 s. Evidently, a larger fraction of stations had intensities exceeding RotD50  
 288 in the transverse orientation when compared to those in the radial orientation. For very short periods of  
 289 vibration, the intensity in the transverse orientation exceeded RotD50 in slightly more than 50% of the  
 290 recording stations, with a tendency for this fraction to increase as the period of vibration increases. For

291 periods of vibrations above 4 s, over 80% of the recording stations had an intensity in the transverse  
292 orientation that exceeds the RotD50 intensity.

293 The extent by which RotD50 intensity is exceeded in the transverse and radial orientation can be  
294 better understood by considering the mean variation of  $S_{aT}/S_{a_{RotD50}}$  and  $S_{aR}/S_{a_{RotD50}}$  with changes in the  
295 period of vibration of the oscillator. Since the intensity at the transverse or radial orientation can take a  
296 maximum value of RotD100, and since the maximum possible value for RotD100/RotD50 is  $\sqrt{2}$  (which  
297 occurs, for example, in a fully polarized record, that is, one in which all the motion occurs at a specific  
298 orientation), both  $S_{aT}/S_{a_{RotD50}}$  and  $S_{aR}/S_{a_{RotD50}}$  range between 0 and  $\sqrt{2}$ . Figure 12 shows the mean ratio  
299 between the ground motion intensity at the transverse or radial orientation and the RotD50 intensity for  $T$   
300  $\in [0.1 \text{ s}, 10 \text{ s}]$ . For oscillators with a short period of vibration between 0.1 s and 0.3 s, the intensity at the  
301 transverse and radial orientations is, on average, approximately equal to RotD50. However, as the oscillator  
302 period increases, the average intensity at the transverse orientation increases from 1.0 to 1.26 times the  
303 RotD50 intensity. In contrast, the radial orientation is characterized by a significant reduction in intensity  
304 relative to the RotD50 intensity with increasing period. Additionally, as previously observed for the  
305 intensities relative to RotD100, the ratio of the intensity at the transverse orientation to RotD50 has a  
306 variability much smaller than that of the  $S_{aR}/S_{a_{RotD50}}$  ratios. The mean ratios shown in Figure 12 could be  
307 used as correction factors to RotD50 intensities estimated using existing GMMs to estimate the level of  
308 intensity in the transverse orientation. However, it is unlikely for a given structure to have primary  
309 orientations aligned with the transverse orientation. Therefore, more importantly, the results motivate the  
310 need to develop modification factors for all orientations so that existing GMMs can be used to estimate  
311 spectral accelerations in any specific orientation with respect to the transverse orientation.

312 To get a better understanding of the influence of epicentral distance on the intensity at the transverse  
313 orientation, both  $S_{aT}/S_{a_{RotD100}}$  and  $S_{aT}/S_{a_{RotD50}}$  were binned by epicentral distance for all stations that  
314 recorded the 2022 event. Figure 13 shows the influence of epicentral distance on histograms of the ratio  
315 between intensity at the transverse orientation and RotD100 for oscillators with periods of 3, 5, 7, and 10  
316 s. Although there are only 12 records available at epicentral distances less than 30 km, it appears that the  
317 intensity at the transverse orientation is more uniformly distributed in this bin. This is consistent with the  
318 findings discussed in Figure 8, where it was shown that the orientation of the maximum intensity with  
319 respect to the transverse orientation was more uniformly distributed in this distance range. In contrast, at  
320 distances greater than 30 km, the distribution of  $S_{aT}/S_{a_{RotD100}}$  is clustered towards the right with the intensity  
321 at the transverse orientation being, on average, 94% to 97% of RotD100. In general, it is apparent that the  
322 intensity at the transverse orientation, on average, gets closer to the RotD100 intensity as epicentral distance  
323 increases, although further studies should be conducted using a larger database of ground motions to better

324 estimate this distance dependence. Figure 14 shows the same histograms binned by distance but for the  
325 ratio between the intensity at the transverse orientation and RotD50. Similar trends are observed where the  
326 intensity at the transverse orientation is, on average, close to the RotD50 intensity in distances less than 30  
327 km, but exceeds RotD50 significantly at longer distances. The implication of this is that RotD50 may not  
328 be the best intensity to use in earthquake-resistant design as it has a significant probability of being exceeded  
329 and that correction factors to existing RotD50 GMMs should be developed to estimate the level of intensity  
330 at specific orientations and that these modification/correction factors may need to account for larger  
331 amplification of RotD50 at larger source-to-site distances.

332 Poulos and Miranda's (2023) main motivation for studying the orientation of the maximum  
333 intensity with respect to the transverse orientation was that S-waves from a theoretical strike-slip double  
334 couple point source with a vertical dip in a homogenous medium have SH radiation pattern that is more  
335 dominant when compared to the SV radiation pattern. Using the epicenter as the source of the radiation  
336 patterns, the orientation of the SH waves coincides with the transverse orientation identified above and in  
337 the previous section. Therefore, the observations made for this earthquake may be explained by the  
338 polarization of S-waves and are therefore primarily the result of a source effect. Site effects such as local  
339 geologic heterogeneities, basin edge effects, and topographic irregularities can also cause polarization  
340 through filtration of response. Regardless, the orientation of the maximum intensity remains close to the  
341 transverse, suggesting the orientation of maximum response is primarily due to source effects.

342 The discussion above and in the prior section focuses mostly on observations for oscillators with  
343 fundamental periods longer than 3 s. While these observations are valuable, it is important to highlight the  
344 applicability of the findings for short-period oscillators as well since they typically have large intensities  
345 and since most structures tend to have first-mode periods in the short-period range (i.e., 0.1 s to 1 s). From  
346 Figures 3, 6, 10, and 11, it is apparent that the observations made for oscillators with periods  $\geq 3$  s also hold  
347 for shorter periods, albeit they are less pronounced. As the oscillator period increases from the short-period  
348 range to the long-period range, the intensity at the transverse orientation significantly increases and the  
349 orientation of the maximum response gets closer to the transverse orientation. The increased alignment of  
350 the orientation of the maximum intensity to the transverse orientation with the increase in oscillator period  
351 may be attributed to wave scattering. More specifically, higher frequencies have shorter wavelengths,  
352 making them more sensitive to heterogeneities in the Earth's crust.

## 353 **5. Summary and Conclusions**

354 This work investigated the orientation and intensity of maximum ground motion intensity for 5%-  
355 damped linear elastic oscillators subjected to the December 20, 2022  $M_w$  6.4 Ferndale earthquake. The  
356 response of oscillators with periods ranging from 0.1 to 10 s was found to be fairly polarized over the

357 geographic area considered. The orientation of the maximum intensity was found to be, in general,  
358 relatively close to the transverse orientation. For example, the fraction of recording stations where the  
359 orientation of the maximum response occurred within  $\pm 25^\circ$  of the transverse orientation was found to be,  
360 depending on the period of vibration, between 2 (~50%) and 3 (~80%) times the fractions that would occur  
361 if the orientation of the maximum response were to be fully random, that is, if all orientations were equally  
362 likely to experience the maximum intensity. On average, the orientation of the maximum intensity gets  
363 closer to the transverse orientation as the period of the oscillator increases. At epicentral distances less than  
364 30 km, the angular distance between the orientation of the maximum intensity and transverse orientation  
365 was found to be more uniformly distributed and without a clear trend for the orientation of the maximum  
366 response. At distances greater than 30 km, the orientation of the maximum intensity is close to the transverse  
367 orientation. These findings are consistent with the recent study by Poulos and Miranda (2023) for strike-  
368 slip earthquakes.

369 Ground motion intensity at the transverse orientation was found to be, on average, close to the  
370 RotD100 intensity for oscillators with periods ranging from 0.1 to 10 s, and these intensities tend to get  
371 closer as oscillator period increases. For example, on average, the intensity at the transverse orientation is  
372 91% of RotD100 intensity at 3 s, and this increases to 95% of RotD100 for 10 s. More importantly, it was  
373 found that the intensity in the transverse orientation is on average only 10% smaller than the maximum  
374 intensity for a wide range of angular differences with respect to the orientation of the maximum intensity  
375 (i.e., the intensity at the transverse orientation was at least 90% of RotD100 up to an  $|\alpha|$  value of  $25^\circ$ ).  
376 Additionally, the intensity at the transverse orientation was found to be systematically larger than RotD50,  
377 with over 80% of the recording stations experiencing intensities in the transverse orientation that exceeded  
378 the RotD50 intensity for periods longer than 4 s. As the oscillator period increases, the average intensity in  
379 the transverse orientation is between 1.0 to 1.26 times the RotD50 intensity. The mean  $S_{aT}/S_{aRotD50}$  ratios  
380 computed in this study could be used as modification/correction factors to RotD50 intensities estimated  
381 with existing GMMs in order to estimate the level of intensity at the transverse orientation, which may be  
382 a good estimator of maximum intensity. Furthermore, the correction factors to RotD50 were found to have  
383 a tendency to increase with increasing source-to-site distance, suggesting that future use of correction  
384 factors to estimate intensity at transverse orientation may need to account for larger amplification of RotD50  
385 with an increase in distance. At source-to-site distances less than 30 km, the intensity at the transverse  
386 orientation is more uniformly distributed, but at distances greater than 30 km the average intensity at the  
387 transverse orientation is 94% to 97% of the RotD100 intensities. Further studies could be carried out to  
388 estimate this dependence on source-to-site more precisely using a larger database of ground motions  
389 recorded from several strike-slip earthquakes.

390 **Author Statement**

391 **Data and Resources**

392 Ground motion records used for the 2022  $M_w$  6.4 Ferndale earthquake were obtained from the  
393 Center for Engineering Strong Motion Data (<https://www.strongmotioncenter.org/>, last accessed February  
394 2023). The network or agencies providing data in this work are the California Strong Motion  
395 Instrumentation Program (CSIMP), the USGS National Strong Motion Project (NSMP), the Berkley Digital  
396 Seismograph Network (BDSN), and the UGS Northern California Regional Network (NCSN). Station  
397 Code/ID of those used in this work are provided below.

398

**List of CESMD Station Code/ID Used in Study and Corresponding Site Characteristics**

Code/ID	Site Class	Code/ID	Site Class	Code/ID	Site Class	Code/ID	Site Class	Code/ID	Site Class
89005	C	1581	N/A	WHMT	N/A	RVIT	N/A	KMPB	N/A
79046	C	1584B	N/A	AONC	N/A	SAGE	N/A	KMR	N/A
79668	C	1023	N/A	BAKR	N/A	TRIN	N/A	KPR	N/A
89255	C	1582	N/A	BJES	N/A	WEAV	N/A	KTD	N/A
89462	D	1586	N/A	BRIC	N/A	KBN	N/A	LASH	N/A
89486	C	1591	N/A	BRIT	N/A	KBU	N/A	LBL	N/A
89641	D	BIGV	N/A	DMOR	N/A	KCO	N/A	LBR	N/A
89688	C	KNEE	N/A	EAGL	N/A	KCR	N/A	LDH	N/A
89781	D	LAB	N/A	GRPK	N/A	KCSB	N/A	LDL	N/A
99640	C	LBE	N/A	HATC	N/A	KCT	N/A	LGY	N/A
99700	C	LBU	N/A	HAYF	N/A	KFP	N/A	LHE	N/A
89734	D	LGMB	N/A	LCOW	N/A	KHBB	N/A	LMC	N/A
89687	D	LWHB	N/A	PETL	N/A	KHMB	N/A	LMI	N/A
1580	N/A	PRDS	N/A	RBOW	N/A	KIP	N/A	LTC	N/A

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405 States Geological Survey (USGS) for installing and maintaining seismic recording stations, and for  
406 collecting, processing, and distributing the ground motion records from the 2022 Ferndale earthquake. This  
407 investigation would not have been possible without these strong motion records.

408 **References**

- 409 Abrahamson, N.A., & Silva, W.J. (1997). Empirical response spectral attenuation relations for shallow  
410 crustal earthquakes. *Seismological Research Letters*, 68(1), 94-127.
- 411 American Society of Civil Engineers. (2010). Minimum Design Loads for Buildings and Other Structures.  
412 *ASCE 7-10 Standard*, American Society of Civil Engineers, Reston, Virginia.
- 413 American Society of Civil Engineers. (2016). Minimum Design Loads for Buildings and Other Structures.  
414 *ASCE 7-16 Standard*, American Society of Civil Engineers, Reston, Virginia.
- 415 American Society of Civil Engineers. (2022). Minimum Design Loads for Buildings and Other Structures.  
416 *ASCE 7-22 Standard*, American Society of Civil Engineers, Reston, Virginia.
- 417 Ancheta, T.D., Darragh, R.B., Stewart, J.P., Seyhan, E., Silva, W.J., Chiou, B.S.J., Wooddell K.E., Graves,  
418 R.W., Kottke, A.R., Boore, D.M., Kishida, T. & Donahue, J.L. (2014). NGA-West2  
419 database. *Earthquake Spectra*, 30(3), 989-1005.
- 420 Bakun, W.H. (2000). Seismicity of California's North Coast. *Bulletin of the Seismological Society of*  
421 *America* 2000; 90 (4): 797–812.
- 422 Beyer, K., & Bommer, J.J. (2006). Relationships between median values and between aleatory variabilities  
423 for different definitions of the horizontal component of motion. *Bulletin of the Seismological Society*  
424 *of America*, 96(4A), 1512-1522.
- 425 Boore, D.M., Joyner, W.B., & Fumal, T.E. (1997). Equations for estimating horizontal response spectra  
426 and peak acceleration from western North American earthquakes: A summary of recent  
427 work. *Seismological Research Letters*, 68(1), 128-153.
- 428 Boore, D.M., & Akkar, S. (2003). Effect of causal and acausal filters on elastic and inelastic response  
429 spectra. *Earthquake Engineering & Structural Dynamics*, 32(11), 1729-1748.
- 430 Boore, D.M. (2004). Choosing the Lowest Usable Frequency for Response Spectra from Filtered Data.  
431 Dave's Notes available at [www.daveboore.com/daves\\_notes](http://www.daveboore.com/daves_notes).
- 432 Boore, D.M., Watson-Lamprey, J., & Abrahamson, N.A. (2006). Orientation-independent measures of  
433 ground motion. *Bulletin of the Seismological Society of America*, 96(4A), 1502-1511.
- 434 Boore, D.M. (2010). Orientation-independent, nongeometric-mean measures of seismic intensity from two  
435 horizontal components of motion. *Bulletin of the Seismological Society of America*, 100(4), 1830-  
436 1835.
- 437 Boore, D.M., Stewart, J. P., Seyhan, E., & Atkinson, G. M. (2014). NGA-West2 equations for predicting  
438 PGA, PGV, and 5% damped PSA for shallow crustal earthquakes. *Earthquake Spectra*, 30(3), 1057-  
439 1085.
- 440 Center for Engineering Strong Motion Data (CESMD). Retrieved on January 18, 2023, from  
441 <https://www.strongmotioncenter.org/index.html>
- 442 Hong, H.P., & Goda, K. (2007). Orientation-dependent ground-motion measure for seismic-hazard  
443 assessment. *Bulletin of the Seismological Society of America*, 97(5), 1525-1538.
- 444 Huang, Y.N., Whittaker, A.S., & Luco, N. (2008). Maximum spectral demands in the near-fault  
445 region. *Earthquake Spectra*, 24(1), 319-341.
- 446 Joyner, W.B., & Boore, D.M. (1982). Estimation of response-spectral values as functions of magnitude,  
447 distance, and site conditions. *Open File Report 82-881*, U.S. Geological Survey.

448 NEHRP Consultants Joint Venture (2011). Selecting and Scaling Earthquake Ground Motions for  
449 Performing Response-History Analysis. *Tech. Rep. NIST GCR 11-917-15*, National Institute of  
450 Standards and Technology, Washington, D.C.

451 Pacific Earthquake Engineering Research Center (PEER) (2017), Guidelines for Performance-Based  
452 Seismic Design of Tall Buildings, Version 2.03, *PEER Report 2017/06*, University of California at  
453 Berkeley, Berkeley, California.

454 Petri, A.E., & Lin, S. (2022, December 21). Northern California towns tally earthquake damage as power  
455 is restored to tens of thousands. *Los Angeles Times*. Retrieved March 05, 2023, from  
456 <https://www.latimes.com/california/story/2022-12-21/california-earthquake-aftermath-damage>

457 Poulos, A., & Miranda, E. (2022a). Proposal of orientation-independent measure of intensity for  
458 earthquake-resistant design. *Earthquake Spectra*, 38(1), 235-253.

459 Poulos, A., Miranda, E., & Baker, J.W. (2022). Evaluation of earthquake response spectra directionality  
460 using stochastic simulations. *Bulletin of the Seismological Society of America*, 112(1), 307-315.

461 Poulos, A., & Miranda, E. (2022b). Probabilistic characterization of the directionality of horizontal  
462 earthquake response spectra. *Earthquake Engineering & Structural Dynamics*, 51(9), 2077-2090.

463 Poulos, A., & Miranda, E. (2023). Effect of faulting type on the orientation of maximum horizontal  
464 earthquake response spectra. *Bulletin of the seismological Society of America*,  
465 doi: <https://doi-org.stanford.idm.oclc.org/10.1785/0120230001>

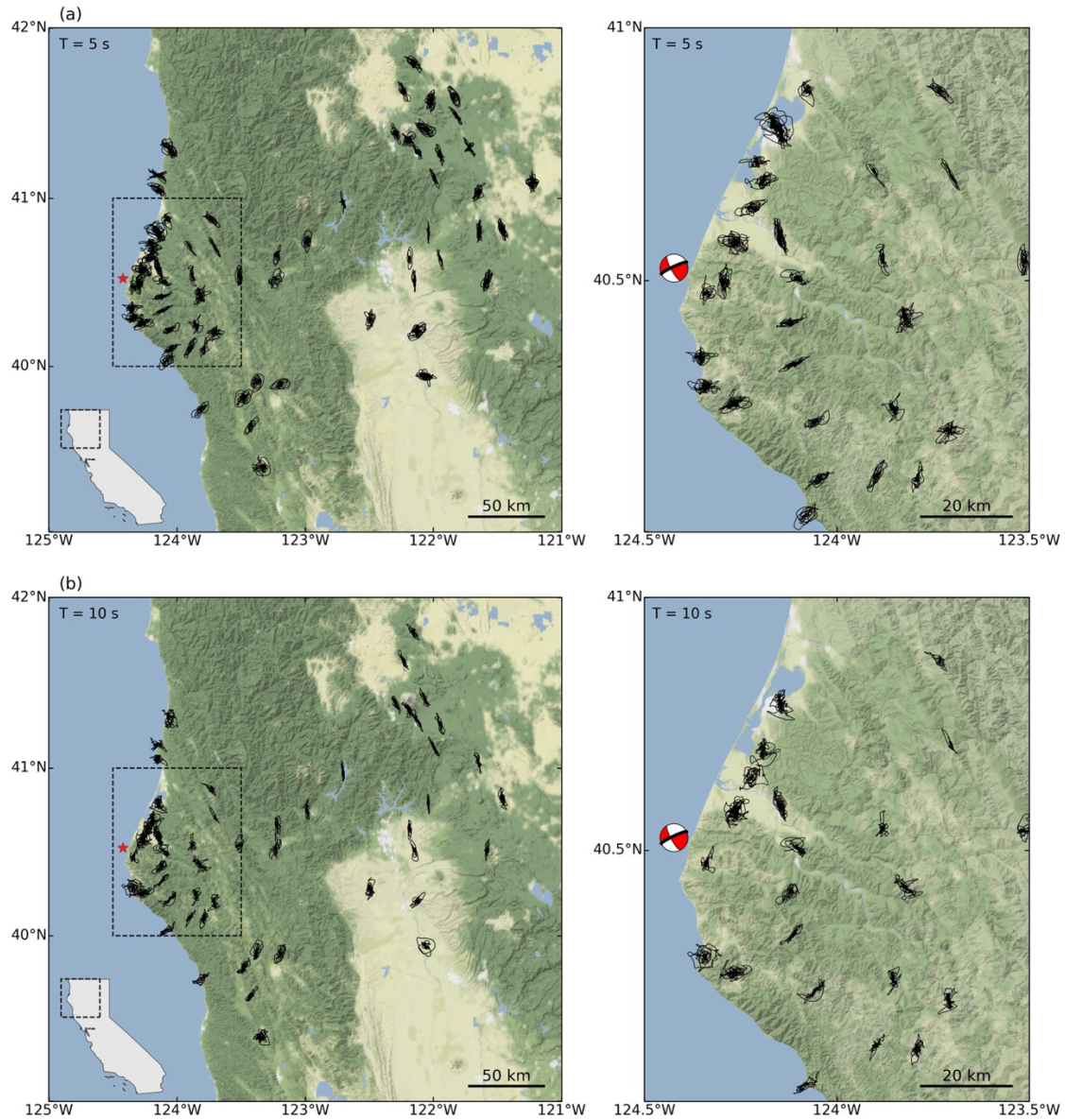
466 Shahi, S.K., & Baker, J.W. (2014). NGA-West2 models for ground motion directionality. *Earthquake*  
467 *Spectra*, 30(3), 1285-1300.

468 Somerville, P.G., Smith, N.F., Graves, R.W., & Abrahamson, N.A. (1997). Modification of empirical strong  
469 ground motion attenuation relations to include the amplitude and duration effects of rupture  
470 directivity. *Seismological Research Letters*, 68(1), 199-222.

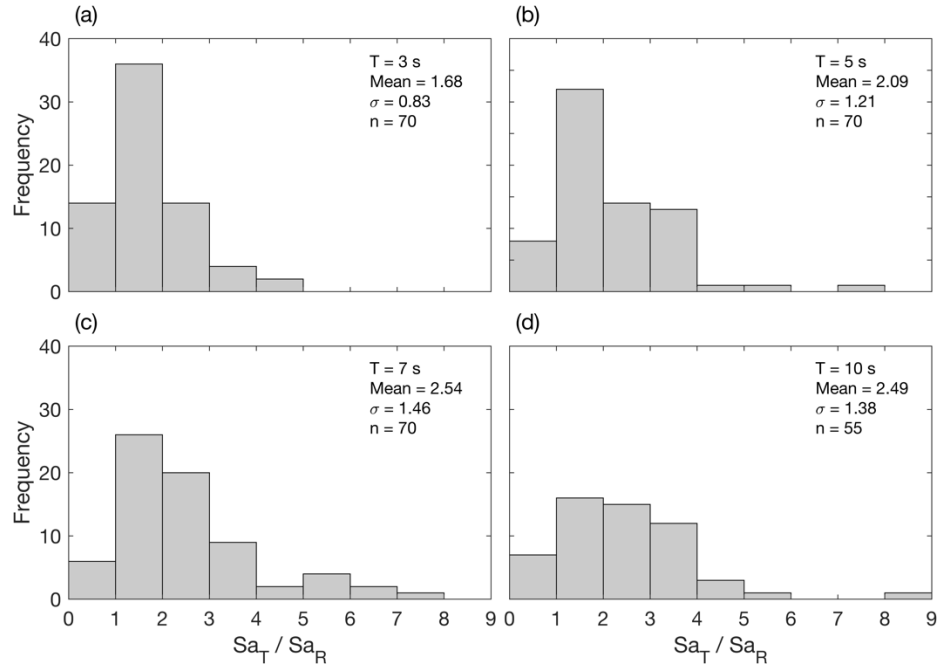
471 Temblor. (2022, December 20). Earthquake rumbles northern California triple junction. Retrieved on  
472 March 05, 2023, from [https://temblor.net/earthquake-insights/dec-20-2022-ferndale-earthquake-](https://temblor.net/earthquake-insights/dec-20-2022-ferndale-earthquake-northern-california-14791/)  
473 [northern-california-14791/](https://temblor.net/earthquake-insights/dec-20-2022-ferndale-earthquake-northern-california-14791/)

474 United States Geological Survey (USGS) (2022). M 6.4 - 15km WSW of Ferndale, CA. Retrieved March  
475 5, 2023, from <https://earthquake.usgs.gov/earthquakes/eventpage/nc73821036/executive>  
476

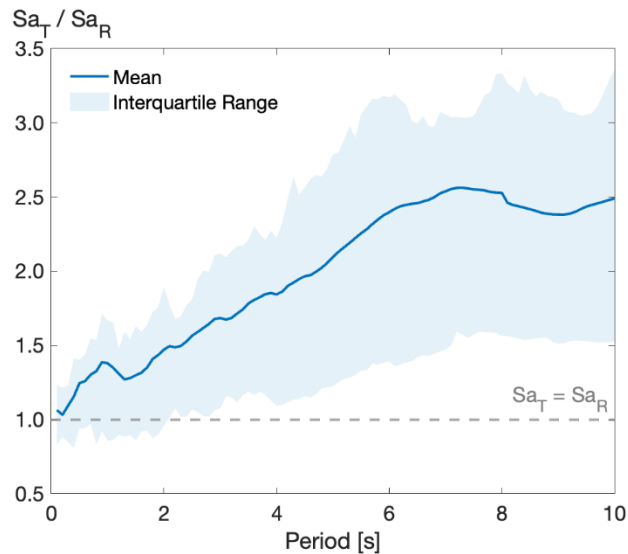




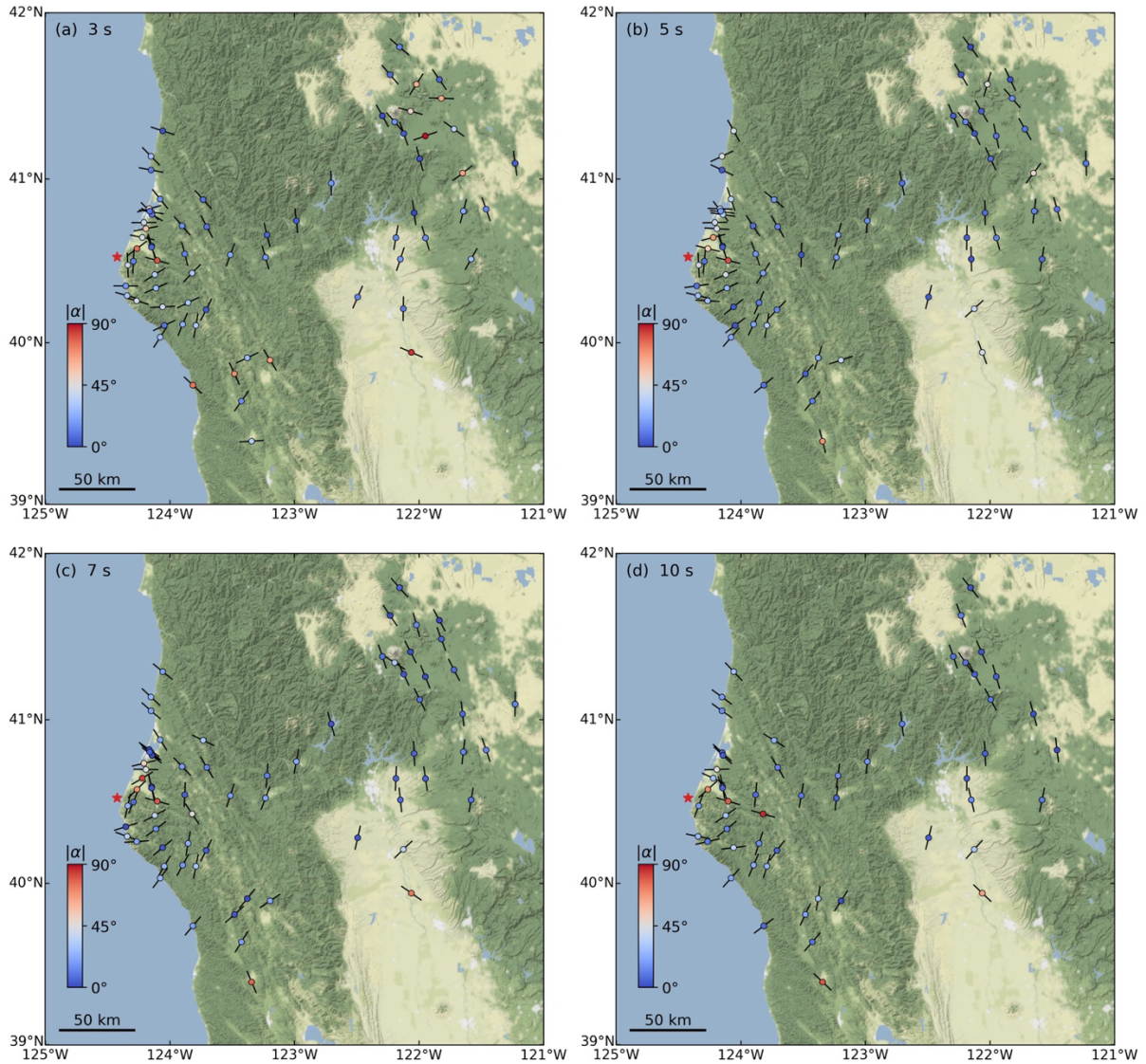
**Figure 1:** Spatial distribution of relative displacement hodographs of 5% damped linear elastic oscillators subjected to recorded ground motions from the 2022  $M_w$  6.4 Ferndale earthquake for (a)  $T = 10$  s, and (b)  $T = 5$  s. The epicenter is indicated by the red star and the thicker black line in the focal mechanism indicates the preferred fault plane.



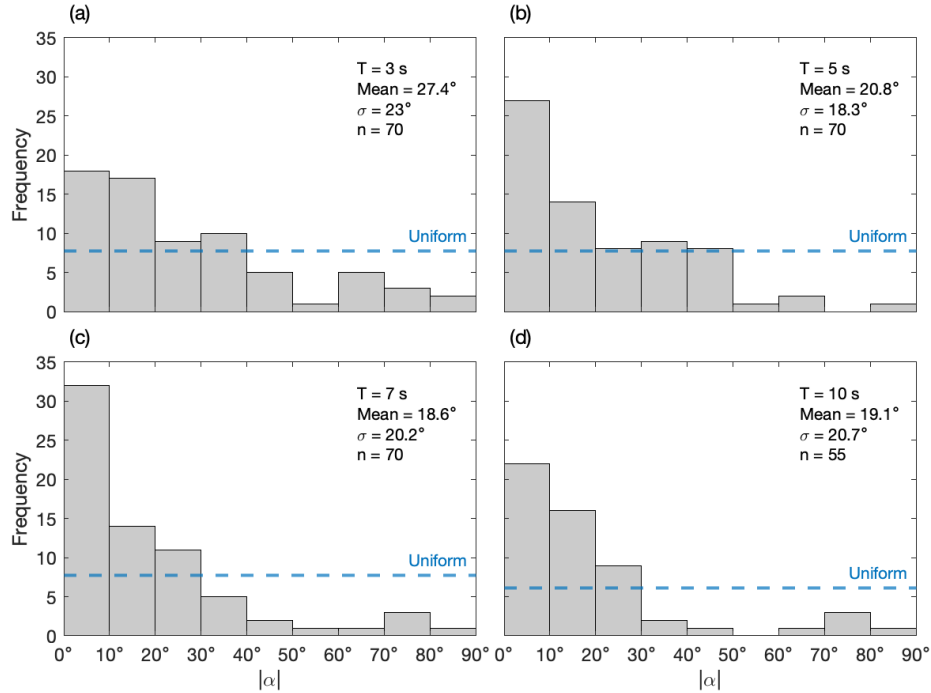
**Figure 2:** Histograms of the ratio between the ground motion intensity in the transverse orientation to the intensity in the radial orientation for (a)  $T = 3$  s, (b)  $T = 5$  s, (c)  $T = 7$  s, and (d)  $T = 10$  s. Each panel also presents the mean and standard deviation ( $\sigma$ ) of the ratio, and the number of records used for each case ( $n$ ).



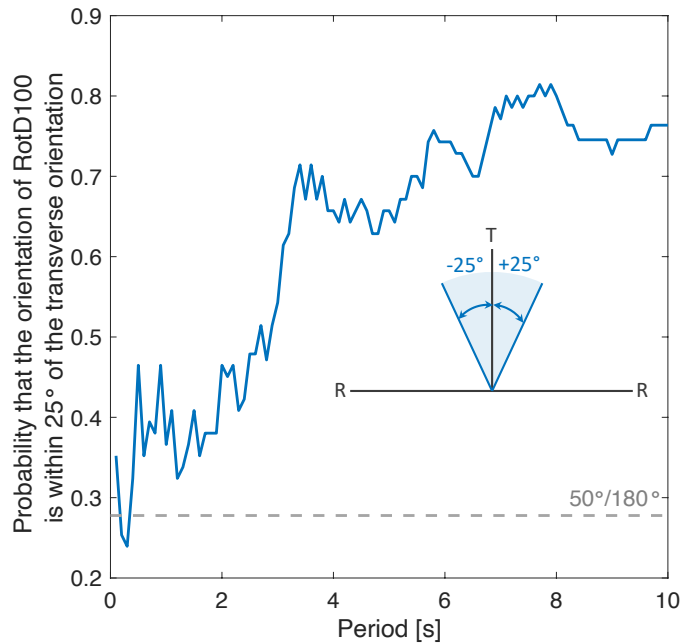
**Figure 3:** Influence of period of vibration on the level of polarization as measured by the mean ratio between the ground motion intensity in the transverse orientation to the intensity in the radial orientation for periods between 0.1 and 10 s.



**Figure 4:** Orientation of the maximum response of 5% damped linear elastic oscillators (indicated by short black lines at each recording station) and their angular distance with respect to the transverse orientation (indicated by the color in each circle) for (a)  $T = 3$  s, (b)  $T = 5$  s, (c)  $T = 7$  s, and (d)  $T = 10$  s.

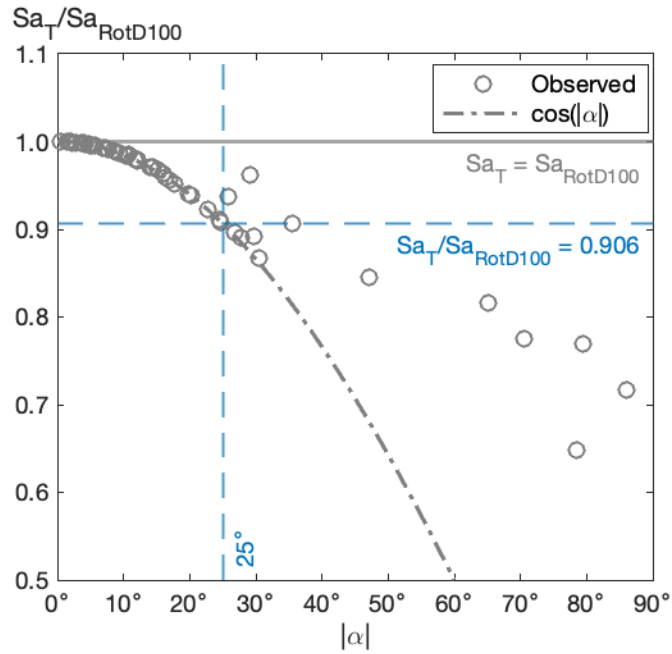


**Figure 5:** Histogram of angular distance between the orientation of RotD100 and the transverse orientation for oscillators with periods of: (a)  $T = 3$  s, (b)  $T = 5$  s, (c)  $T = 7$  s, and (d)  $T = 10$  s. Dotted line represents the histogram if the orientation of RotD100 were to be equally likely in all orientations. Each panel also presents the mean and standard deviation ( $\sigma$ ) of the angles, and the number of records used for each case ( $n$ ).

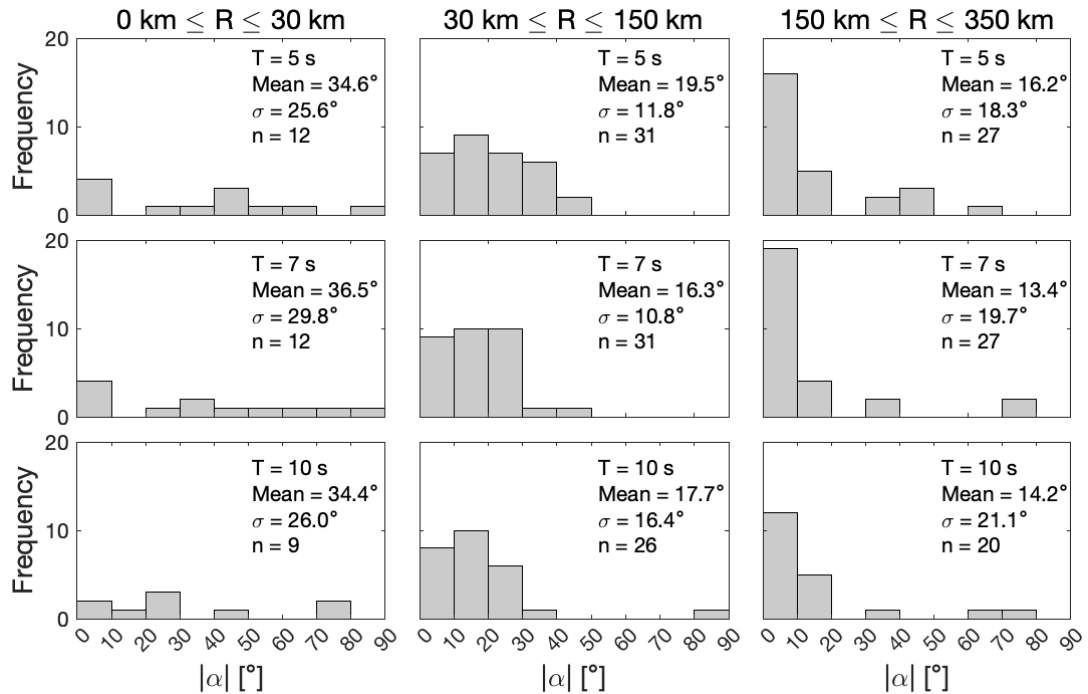


**Figure 6:** Fraction of stations where the orientation of RotD100 falls within  $\pm 25^\circ$  of the transverse orientation. For a fully random case, the expected fraction would be  $50^\circ/180^\circ$  (0.28).

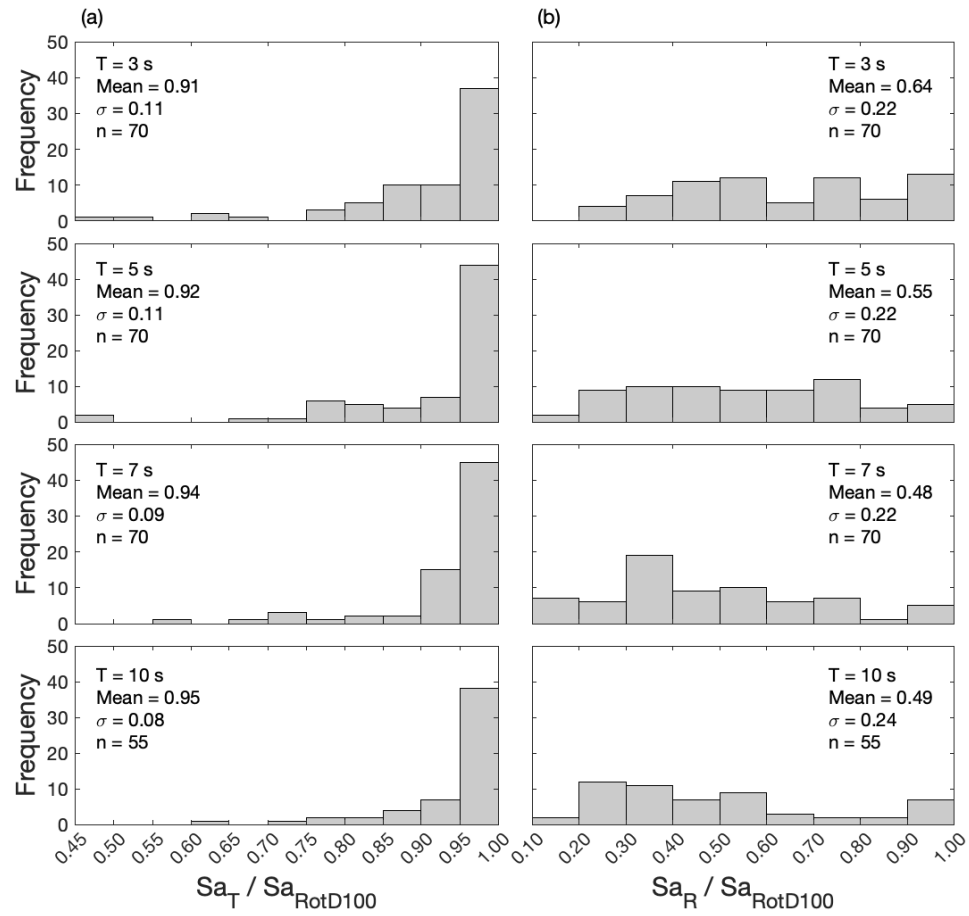




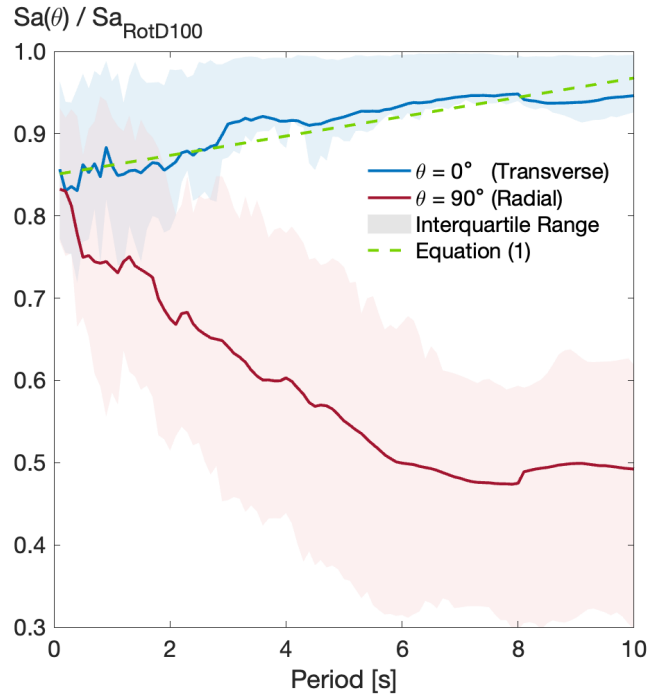
**Figure 7:** Relationship between  $|\alpha|$  and the ratio of the intensity at the transverse orientation to the maximum intensity at each recorded station for a period of 10 s.



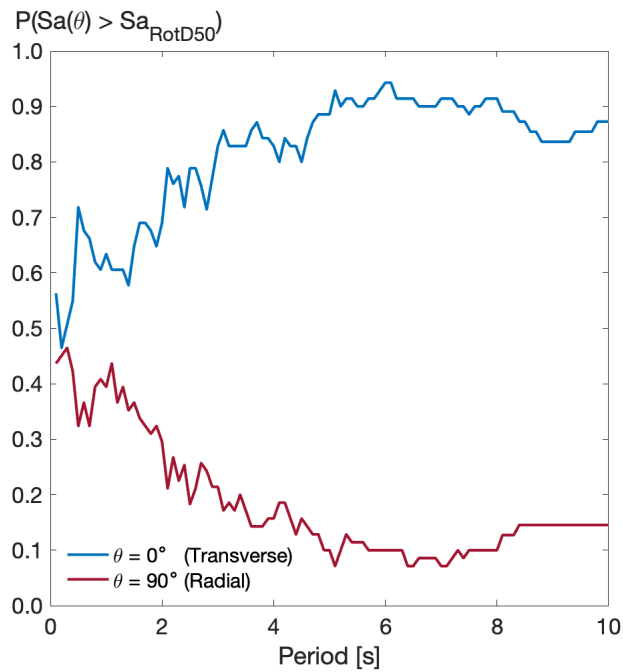
**Figure 8:** Influence of epicentral distance ( $R$ ) on the histograms of angular difference between the orientation of RotD100 and the transverse orientation for oscillators with periods of 5, 7, and 10 s. Each panel presents the mean and standard deviation ( $\sigma$ ) of the angles, and the number of records used for each case ( $n$ ).



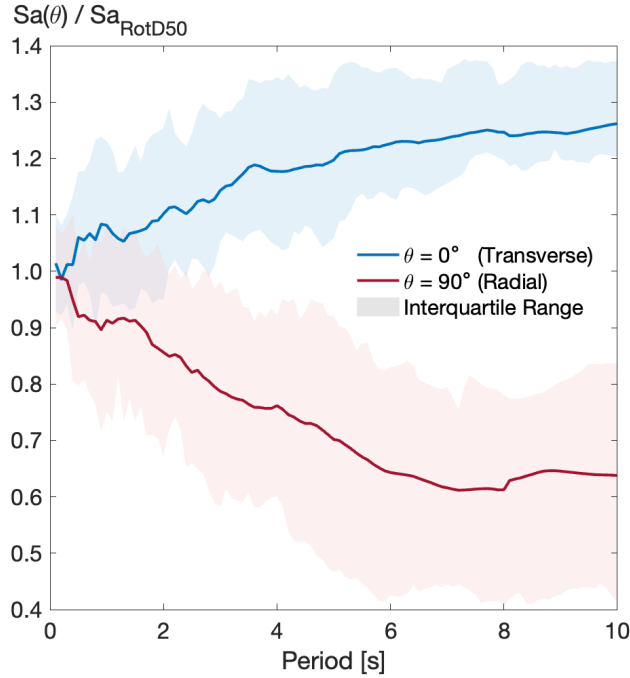
**Figure 9:** Histogram of the ratio between the ground motion intensity at the (a) transverse or (b) radial orientation and RotD100 for periods of 3, 5, 7, and 10 s. Each panel presents the mean and standard deviation ( $\sigma$ ) of the ratios, and the number of records used for each case ( $n$ ).



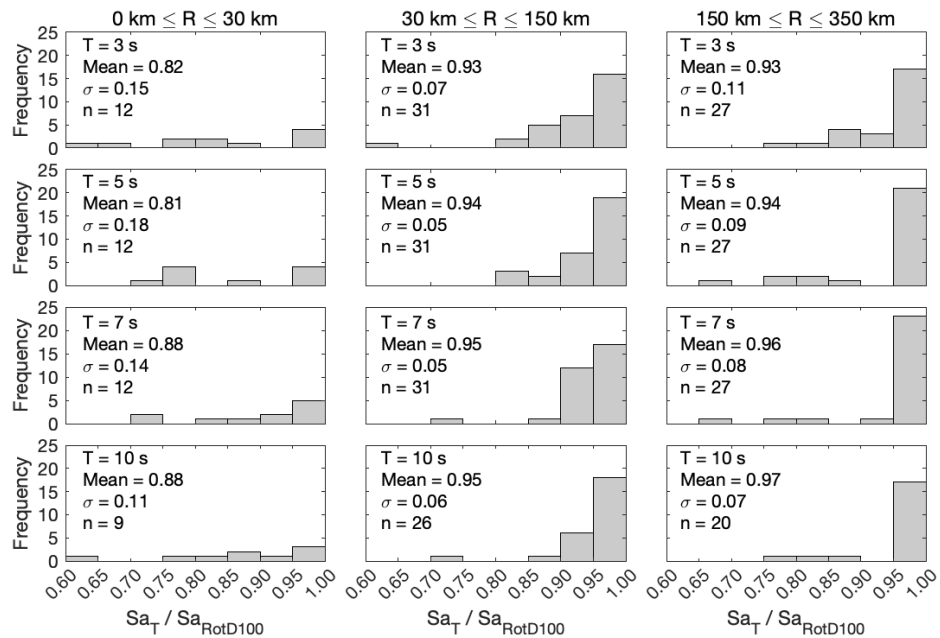
**Figure 10:** Mean ratio between the ground motion intensity at the transverse orientation and the RotD100 intensity and between the intensity at the radial orientation and the RotD100 intensity for periods between 0.1 and 10 s. Shaded bands around means show the interquartile ranges at each period.



**Figure 11:** Fraction of stations where the intensity at the transverse or radial orientation exceeds the RotD50 intensity.

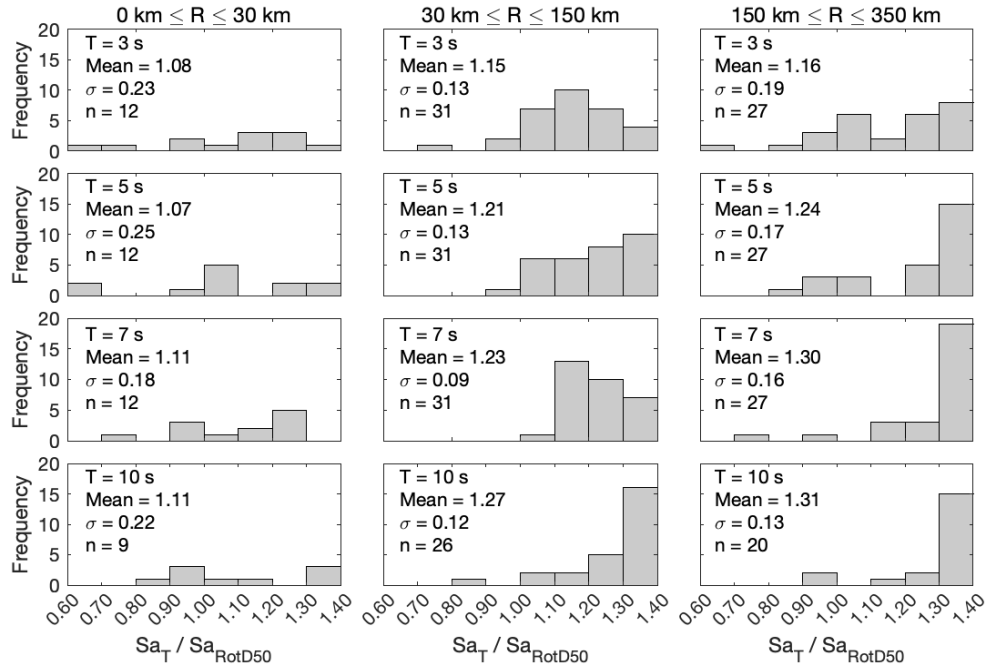


**Figure 12:** Mean ratio between the ground motion intensity at the transverse orientation and the RotD50 intensity and between the intensity at the radial orientation and the RotD50 intensity for periods between 0.1 and 10 s. Shaded bands around means show the interquartile ranges at each period.



**Figure 13:** Influence of epicentral distance (R) on histograms of the ratio between the ground motion intensity at the transverse orientation and RotD100 for oscillators with periods of 3, 5, 7, and 10 s. Each panel presents the mean and standard deviation ( $\sigma$ ) of the ratios, and the number of records used for each case (n).





**Figure 14:** Influence of epicentral distance ( $R$ ) on histograms of the ratio between the ground motion intensity at the transverse orientation and RotD50 for oscillators with periods of 3, 5, 7, and 10 s. Each panel presents the mean and standard deviation ( $\sigma$ ) of the ratios, and the number of records used for each case ( $n$ ).